# Lead-tellurium oxysalts from Otto Mountain near Baker, California: III. Thorneite, $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right) \mathrm{Cl}_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)$, the first mineral with edge-sharing octahedral tellurate dimers 

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#### Abstract

Thorneite, $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right) \mathrm{Cl}_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)$, is a new tellurate from Otto Mountain near Baker, California, named in honor of Brent Thorne. The new mineral occurs on fracture surfaces and in small vugs in brecciated quartz veins. Thorneite is directly associated with acanthite, cerussite, gold, hessite, iodargyrite, khinite, wulfenite, and three other new tellurates: housleyite, markcooperite, and ottoite. Various other secondary minerals occur in the veins, including three other new secondary tellurium minerals: paratimroseite, telluroperite, and timroseite. Thorneite is monoclinic, space group $C 2 / c, a$ $=21.305(1), b=11.059(1), c=7.564(1) \AA, \beta=101.112(4)^{\circ}, V=1748.8(4) \AA^{3}$, and $Z=4$. Crystals are prismatic to bladed with elongation and striations parallel to $\mathbf{c}$ and typically occur in parallel and random aggregates. It is yellow and transparent, with pale yellow streak and adamantine luster. Mohs hardness is estimated at 2 . The mineral is brittle, with an irregular to splintery fracture and good $\{100\}$ cleavage. The calculated density is $6.828 \mathrm{~g} / \mathrm{cm}^{3}$. Thorneite is biaxial ( + ), with large $2 V$, but indices of refraction are too high to be measured. The optic orientation is $Y=\mathbf{b}, Z^{\wedge} \mathbf{a}=29^{\circ}$ in obtuse $\beta$. No pleochroism was observed. Electron microprobe analysis provided $\mathrm{PbO} 73.90, \mathrm{ZnO} 0.03, \mathrm{TeO}_{3} 20.35$, $\mathrm{Cl} 2.29, \mathrm{H}_{2} \mathrm{O} 1.28$ (structure), $\mathrm{CO}_{2} 2.29$ (structure), $\mathrm{O} \equiv \mathrm{Cl}-0.52$, total $99.62 \mathrm{wt} \%$; the empirical formula (based on $\mathrm{O}+\mathrm{Cl}=16$ ) is $\left(\mathrm{Pb}_{5.94} \mathrm{Zn}_{0.01}\right)\left(\mathrm{Te}_{2.08}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{C}_{1.00} \mathrm{O}_{3}\right)\left[\mathrm{Cl}_{1.16} \mathrm{O}_{0.34}(\mathrm{OH})_{0.50}\right]\left(\mathrm{H}_{2} \mathrm{O}\right)$. The strongest powder X-ray diffraction lines are [ $d_{\text {obs }}$ in $\left.\AA(h k l) I\right]: 10.43(200) 35,3.733(\overline{5} 11, \overline{2} 02,002) 27,3.595$ ( $\overline{4} 21$ ) $33,3.351(112) 66,3.224(511,131) 100,3.093(\overline{2} 22, \overline{3} 31) 30,2.900(\overline{6} 21) 44,2.133(821,622$, $223,731,242) 38$. The crystal structure $\left(R_{1}=0.028\right)$ contains edge-sharing octahedral tellurate dimers, $\left[\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right]^{8-}$ that bond to Pb atoms, which in turn are linked via bonds to Cl atoms, $\mathrm{CO}_{3}$ triangles, and $\mathrm{H}_{2} \mathrm{O}$ molecules.


Keywords: Thorneite, new mineral, tellurate, crystal structure, Otto Mountain, California

## INTRODUCTION

Thorneite, $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right) \mathrm{Cl}_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)$, is one of seven new secondary lead-tellurium minerals discovered recently at Otto Mountain near Baker, California. Detailed information on the mining history, geology, mineralogy, and mineral paragenesis of the deposit, as well as the discovery of the new minerals, is provided in Kampf et al. (2010a).

Thorneite is named in honor of Brent Thorne (b. 1951) of Bountiful, Utah, the discoverer of the mineral. Brent Thorne is an avid field collector and regular contributor to Mindat.org with over 1100 photographs posted on the Mindat website. He has collected and provided specimens for the description of the new mineral plumbophyllite (Kampf et al. 2009) and for continuing studies of other new and existing minerals from Otto Mountain and the nearby Blue Bell claims. Thorne has agreed to the naming of the mineral in his honor. The new mineral and name have been approved by the Commission on New Minerals, Nomenclature and Classification of the International Mineralogical Association (IMA 2009-023). Four cotype specimens are deposited in

[^0]the Natural History Museum of Los Angeles County, catalog numbers $62257,62258,62259$, and 62260.

## OCCURRENCE

Thorneite was found in the Bird Nest drift on the southwest flank of Otto Mountain, approximately 2 km northwest of Baker, San Bernardino County, California. The Bird Nest drift ( $35^{\circ}$ $16.606^{\prime} \mathrm{N}, 116^{\circ} 05.956^{\prime} \mathrm{W}$ ) is located 0.7 km northwest of the Aga mine ( $35^{\circ} 16.399^{\prime} \mathrm{N}, 116^{\circ} 05.665^{\prime} \mathrm{W}$ ).

Thorneite is rare and occurs mostly on fracture surfaces and in small vugs in quartz veins. Species observed in direct association with the new mineral include acanthite, cerussite, gold, hessite, iodargyrite, khinite, wulfenite, and three other new tellurates: housleyite $\left[\mathrm{Pb}_{6} \mathrm{Cu}^{2+} \mathrm{Te}_{4}^{6+} \mathrm{O}_{18}(\mathrm{OH})_{2}\right]$ (IMA2009-024; Kampf et al. 2010b), markcooperite $\left[\mathrm{Pb}_{2}\left(\mathrm{UO}_{2}\right) \mathrm{Te}^{6+} \mathrm{O}_{6}\right]$ (IMA2009-045; Kampf et al. 2010c), and ottoite $\left[\mathrm{Pb}_{2} \mathrm{Te}^{6+} \mathrm{O}_{5}\right]$ (IMA2009-063; Kampf et al. 2010a). Other species identified in the Otto Mountain assemblages include: anglesite, atacamite, boleite, brochantite, burckhardtite, calcite, caledonite, celestine, chalcopyrite, bromine-rich chlorargyrite, chrysocolla, devilline, diaboleite, fluorite, fornacite, galena, goethite, jarosite, kuranakhite, linarite, malachite, mimetite, mottramite, munakataite, murdochite,
muscovite, perite, phosphohedyphane, plumbojarosite, pyrite, schieffelinite, vanadinite, vauquelinite, and three other new minerals: paratimroseite $\left[\mathrm{Pb}_{2} \mathrm{Cu}_{4}^{2+}\left(\mathrm{Te}^{6+} \mathrm{O}_{6}\right)_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)_{2}\right]$ (IMA2009-065; Kampf et al. 2010d), telluroperite $\left[\mathrm{Pb}_{3} \mathrm{Te}^{4+} \mathrm{O}_{4} \mathrm{Cl}_{2}\right]$ (IMA2009-044; Kampf et al. 2010e), and timroseite $\left[\mathrm{Pb}_{2} \mathrm{Cu}_{5}^{2+}\left(\mathrm{Te}^{6+} \mathrm{O}_{6}\right)_{2}(\mathrm{OH})_{2}\right]$ (IMA2009-064; Kampf et al. 2010d). Other potentially new species are still under investigation.

Thorneite and most of the other secondary minerals of the quartz veins are interpreted as having formed from the partial oxidation of primary sulfides (e.g., galena) and tellurides (e.g., hessite) during or following brecciation of the quartz veins.

## PHYSICAL AND OPTICAL PROPERTIES

The mineral occurs as prismatic to bladed crystals (Fig. 1) with elongation and striations parallel to $\mathbf{c}$. They typically occur in parallel and random intergrowths and are up to 0.5 mm in length. Crystals exhibit the forms $\{100\},\{010\},\{110\}$, and $\{21 \overline{1}\}$ (Fig. 2). Crystal intergrowths are common, but no systematic crystallographic orientation was observed.

Thorneite is yellow and transparent, with pale yellow streak and adamantine luster. The mineral is non-fluorescent. The Mohs hardness is estimated at 2 based upon behavior of crystals when broken. The tenacity is brittle and the fracture is irregular to splintery. Crystals are very fragile and easily crushed. Cleavage is good on $\{100\}$. The density could not be measured because it is greater than those of available high-density liquids and there is insufficient material for physical measurement. The calculated density is $6.790 \mathrm{~g} / \mathrm{cm}^{3}$ for the empirical formula and $6.829 \mathrm{~g} / \mathrm{cm}^{3}$ for formula based on refined occupancies in structure (see below). Thorneite turns opaque and dissolves slowly in HCl .


Figure 1. Backscatter SEM image of thorneite.

The indices of refraction exceed those of available index fluids. The Gladstone-Dale relationship (Mandarino 1981) predicts $n_{\mathrm{av}}=2.021$ based on the empirical formula. Birefringence is very low ( $\sim 0.01$ ), suggesting that all indices of refraction are in the 2.01 to 2.03 range. Orthoscopic and conoscopic optical examination using a Leitz Ortholux I polarizing microscope equipped with a Supper spindle stage showed thorneite to be biaxial $(+)$, with large $2 V$. No dispersion was observed. The optical orientation is $Y=\mathbf{b}, Z^{\wedge} \mathbf{a}=29^{\circ}$ in obtuse $\beta$. The mineral is non-pleochroic.

## RAMAN SPECTROSCOPY

Raman spectroscopic micro-analysis was carried out using a Renishaw M1000 micro-Raman spectrometer system. Spectra were obtained using a 514.5 nm laser with a $100 \times$ objective producing a spot about $2 \mu \mathrm{~m}$ in diameter with about 5 mW power. Peak positions were calibrated against a silicon standard. Spectra were obtained with a dual-wedge polarization scrambler inserted directly above the objective lens to minimize the effects of polarization. Raman measurements were made on an isolated 150 $\mu \mathrm{m} \times 50 \mu \mathrm{~m}$ thorneite crystal on quartz matrix. Spectra obtained with the light propagating parallel and perpendicular to the $\mathbf{c}$ axis of the crystal were found to be very similar. The spectra with the light propagating parallel to the $\mathbf{c}$ axis are shown in Figure 3a (100 to $1600 \mathrm{~cm}^{-1}$ range) and Figure $3 b$ ( 0 to $4000 \mathrm{~cm}^{-1}$ range with the vertical scale expanded).

The sharp peak at $1056 \mathrm{~cm}^{-1}$ is in the shift range where the strongest carbonate peak typically occurs. The broad peak at about $3300 \mathrm{~cm}^{-1}$ is attributable to OH stretching modes and the peak at about $1630 \mathrm{~cm}^{-1}$ is attributable to the HOH bending mode of water. The Raman spectroscopy, therefore, corroborates the presence of carbonate and water indicated by the structure determination (see below).


Figure 2. Crystal drawing of thorneite (clinographic projection).


Figure 3. Raman spectra for thorneite (a) covering the range 100 to $1600 \mathrm{~cm}^{-1}$ and (b) covering the range 0 to $4000 \mathrm{~cm}^{-1}$ with expanded vertical scale.

## Chemistry

Three chemical analyses were carried out using a JEOL8200 electron microprobe (WDS mode, $15 \mathrm{kV}, 10 \mathrm{nA}$, focused beam) at the Division of Geological and Planetary Sciences, California Institute of Technology. The $\mathrm{H}_{2} \mathrm{O}$ and $\mathrm{CO}_{2}$ were calculated by stoichiometry from the results of the crystal-structure analysis. The averages (and ranges) of the analyses are: PbO 73.90 (73.74-73.98), ZnO 0.03 ( $0.02-0.05$ ), $\mathrm{TeO}_{3} 20.35$ (20.05-20.76), Cl 2.29 (1.75-2.98), $\mathrm{H}_{2} \mathrm{O} 1.26$ (structure), $\mathrm{CO}_{2} 2.45$ (structure), $\mathrm{O} \equiv \mathrm{Cl}-0.52$, total $99.76 \mathrm{wt} \%$. Standards used were $\mathrm{PbS}, \mathrm{ZnO}$, Te metal, and sodalite (for Cl ). There was some evidence of beam damage during the EMP analyses; however, the final totals are close to $100 \%$ and the empirical formula derived from the analyses fits the structural formula well.

The empirical formula (based on $\mathrm{O}+\mathrm{Cl}=16$ ) is $\left(\mathrm{Pb}_{5.94} \mathrm{Zn}_{0.01}\right)$ $\left(\mathrm{Te}_{2.08}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{C}_{1.00} \mathrm{O}_{3}\right)\left(\mathrm{Cl}_{1.16} \mathrm{O}_{0.34}(\mathrm{OH})_{0.50}\right)\left(\mathrm{H}_{2} \mathrm{O}\right)$. The simplified formula is $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right)\left(\mathrm{Cl}_{1.50}(\mathrm{OH})_{0.50}\right)\left(\mathrm{H}_{2} \mathrm{O}\right)$, based on the structure with refined Cl and OH occupancies for the $\mathrm{Cl} / \mathrm{OH}$ site. The end-member formula is $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right) \mathrm{Cl}_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)$, which requires $\mathrm{PbO} 74.10, \mathrm{TeO}_{3} 19.43, \mathrm{CO}_{2} 2.44, \mathrm{H}_{2} \mathrm{O} 1.00, \mathrm{Cl} 3.92$, $\mathrm{O}=\mathrm{Cl}-0.89$, total $100.00 \mathrm{wt} \%$.

## X-RAY CRYSTALLOGRAPHY AND STRUCTURE detervination

Both powder and single-crystal X-ray diffraction data were obtained on a Rigaku R-Axis Spider curved imaging plate microdiffractometer utilizing monochromatized Mo $K \alpha$ radiation. The powder data presented in Table 1 show good agreement with the pattern calculated from the structure determination.

Table 1. X-ray powder-diffraction data for thorneite

| $l_{\text {obs }}$ | $d_{\text {obs }}$ | $d_{\text {calc }}$ | $I_{\text {calc }}$ | hkl |
| :---: | :---: | :---: | :---: | :---: |
| 35 | 10.430 | 10.452 | 60 | $200^{*}$ |
| 14 | 9.740 | 9.775 | 25 | 110 |
| 6 | 5.550 | 5.529 | 14 | 020 * |
| 10 | 5.040 | 5.033 | 12 | $\overline{3} 1{ }^{*}$ |
| 13 | 4.940 | 4.888 | 12 | 220 |
| 6 | 4.440 | 4.434 | 9 | 02 1* |
| 18 | 4.270 | 4.288 | 14 | 31 1* |
| 10 | 3.900 | 3.911 | 16 | $510^{*}$ |
| 27 | 3.7333.595 | 3.797 | 11 | 420 |
|  |  | 3.750 | 11 | 511 |
|  |  | 3.731 | 18 | $\overline{2} 02^{*}$ |
|  |  | 3.711 | 13 | 002 |
| 33 |  | 3.591 | 47 | $\overline{4} 21^{*}$ |
|  |  | 3.400 | 15 | $\overline{3} 12$ |
| 66 | 3.351 | 3.370 | 38 | $112^{*}$ |
|  |  | 3.305 | 21 | 131 |
|  |  | 3.302 | 14 | 202 |
| 100 | 3.224 | 3.228 | 73 | $51{ }^{*}$ |
|  |  | 3.219 | 100 | 13 1* |
| 30 | 3.093 | 3.093 | 26 | $\overline{2} 22^{*}$ |
|  |  | 3.088 | 21 | $331 *$ |
| 21 | 2.969 | 2.973 | 31 | 512 |
|  |  | 2.932 | 10 | 312 |
| 44 | 2.900 | 2.907 | 55 | $\overline{6} 21^{*}$ |
|  |  | 2.883 | 18 | 710 |
|  |  | 2.874 | 23 | 711 |
|  |  | 2.835 | 11 | 222 |
| 6 | 2.691 | 2.706 | 7 | 531 |
|  |  | 2.673 | 10 | 240 |
| 8 | 2.606 | 2.613 | 5 | 800 |
|  |  | 2.598 | 4 | 621 |
| 6 | 2.581 | 2.591 | 7 | 041 |
|  |  | 2.444 | 8 | 440 |
| 12 | 2.428 | 2.427 | 18 | $\overline{3} 13^{*}$ |
| 8 | 2.361 | 2.363 | 4 | 820 |
|  |  | 2.362 | 12 | - 0 2* |
| 3 | 2.285 | 2.292 | 4 | 223 |
| 6 | 2.226 | 2.221 | 6 | $\underline{2} 42$ |
|  |  | 2.220 | 7 | 423 |
| 26 | 2.153 | 2.159 | 13 | 313 |
|  |  | 2.146 | 11 | 821 |
|  |  | 2.144 | 16 | 622 |
| 38 | 2.133 | 2.131 | 11 | 223 |
|  |  | 2.126 | 20 | 731 |
|  |  | 2.120 | 14 | 242 |
| 4 | 2.075 | 2.078 | 5 | $\overline{7} 13$ |
|  |  | 2.060 | 5 | $\overline{3} 51$ |
|  |  | 1.973 | 16 | $533^{*}$ |
| 15 | 1.969 | 1.966 | 8 | 802 |
|  |  | 1.955 | 12 | 550 |
|  |  | 1.875 | 5 | 1022 |
|  |  | 1.873 | 9 | $1110^{*}$ |
|  |  | 1.866 | 9 | $\overline{4} 04^{*}$ |
| 15 | 1.863 | 1.858 | 7 | $\overline{9} 13$ |
|  |  | 1.843 | 12 | 060 |
|  |  | 1.798 | 10 | 552 |
|  |  | 1.789 | 11 | 224 |
|  |  | 1.780 | 9 | 642 |
| 22 | 1.782 | 1.777 | 17 | 750 * |
|  |  | 1.773 | 9 | 243 |
|  | 1.737 | 1.742 | 8 | 1200 |
| 3 |  | 1.737 | 4 | 533 |
| 4 | 1.684 | 1.693 | 7 | 314 |
| 6 | 1.665 | 1.665 | 6 | 443 |
| 6 | 1.617 | 1.617 | 7 | $\overline{10} 42^{*}$ |
|  |  | 1.614 | 4 | 1022 |
| 5 | 1.574 | 1.579 | 4 | 733 |
|  |  | 1.561 | 3 | 1221 |
| 10 | 1.555 | 1.561 | 9 | $\underline{2} 44$ |
|  |  | 1.547 | 3 | 952 |
| 12 | 1.493 | 1.497 | 7 | 1331 * |
|  |  | 1.496 | 5 | 371 |

Notes: $I_{\text {obs }}$ based upon peak heights. $I_{\text {calc }}$ calculated from the crystal structure using Powder Cell (Kraus and Nolze 1996). $d_{\text {calc }}$ based on the cell refined from the powder data (*) using UnitCell (Holland and Redfern 1997). Refined cell: $a=21.226(3), b=11.106(2), c=7.548(1)$ $\AA$ A $, \beta=101.33(2)^{\circ}, V=1744.6(4) \AA^{3}$.

The Rigaku CrystalClear software package was used for processing the structure data, including the application of an empirical absorption correction. The structure was solved by direct methods using SIR92 (Altomare et al. 1994) and the location of all non-hydrogen atoms was straightforward. The SHELXL-97 software (Sheldrick 2008) was used, with neutral atom scattering factors, for the refinement of the structure. Bond-valence calculations indicate that one O atom (designated OW ) is a water molecule. A difference Fourier map revealed the H atom site corresponding to both H atoms of the water molecule. The position of the H site was constrained to an OW-H distance of $0.90(3) \AA$ and an $\mathrm{H}-\mathrm{H}$ distance of $1.45(3) \AA$ in the final refinement. The structure refinement indicated the Cl site to be jointly occupied by Cl and $\mathrm{OH}\left[\mathrm{Cl}_{0.75}(\mathrm{OH})_{0.25}\right]$. It was not possible to locate the 0.25 H atom associated with this site.

The details of the data collection and the final structure refinement are provided in Table 2. The final atomic coordinates and displacement parameters are in Table 3. Selected interatomic

TABLE 2. Data collection and structure refinement details for thorneite

| Diffractometer | Rigaku R-Axis Spider |
| :---: | :---: |
| X-ray radiation/power | MoK $\alpha(\lambda=0.71075 \AA) / 50 \mathrm{kV}, 40 \mathrm{~mA}$ |
| Temperature | 298(2) K |
| Structural formula | $\mathrm{Pb}_{6}\left(\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right)\left(\mathrm{CO}_{3}\right)\left(\mathrm{Cl}_{1.50}(\mathrm{OH})_{0.50}\right)\left(\mathrm{H}_{2} \mathrm{O}\right)$ |
| Space group | C2/c |
| Unit-cell dimensions | $a=21.305(1) \AA$ |
|  | $b=11.059(1) \AA$ |
|  | $c=7.564(1) \AA$ |
|  | $\beta=101.112(4)^{\circ}$ |
| Z | 4 |
| Volume | 1748.8(4) $\AA^{3}$ |
| Density (for above formula) | $6.829 \mathrm{~g} / \mathrm{cm}^{3}$ |
| Absorption coefficient | $61.129 \mathrm{~mm}^{-1}$ |
| F(000) | 3016 |
| Crystal size | $85 \times 70 \times 15 \mu \mathrm{~m}$ |
| $\theta$ range | 3.46 to $23.24^{\circ}$ |
| Index ranges | $-23 \leq h \leq 23,-12 \leq k \leq 12,-7 \leq I \leq 8$ |
| Reflections collected/unique | 8170/1255 [ $\left.\mathrm{R}_{\text {int }}=0.0610\right]$ |
| Reflections with $F_{0}>4 \sigma F$ | 1135 |
| Completeness to $\theta=27.50^{\circ}$ | 99.7\% |
| Max. and min. transmission | 0.4608 and 0.0782 |
| Refinement method | Full-matrix least-squares on $F^{2}$ |
| Parameters refined | 120 |
| GoF | 1.069 |
| Final $R$ indices [ $F_{0}>4 \sigma$ F] | $R_{1}=0.028, \mathrm{wR}_{2}=0.054$ |
| $R$ indices (all data) | $R_{1}=0.033, \mathrm{w} R_{2}=0.055$ |
| Largest diff. peak/hole | 1.395/-1.327e/Å ${ }^{3}$ |
| Notes: $R_{\text {int }}=\Sigma \mid F_{o}^{2}-F_{0}^{2}($ mean $)\left\|/ \Sigma\left[F_{0}^{2}\right] . G o F=S=\left\{\Sigma\left[w\left(F_{o}^{2}-F_{\mathrm{c}}^{2}\right)^{2}\right] /(n-p)\right\}^{1 / 2} . R_{1}=\Sigma\right\|\left\|F_{0}\right\|$ $-\left\|F_{\mathrm{c}}\right\|\|\Sigma\| F_{\mathrm{o}} \mid \cdot w R_{2}=\left\{\Sigma\left[w\left(F_{\mathrm{o}}^{2}-F_{\mathrm{c}}^{2}\right)^{2}\right] / \Sigma\left[w\left(F_{0}^{2}\right)^{2}\right]\right\}^{1 / 2} \cdot w=1 /\left[\sigma^{2}\left(F_{o}^{2}\right)+(a P)^{2}+b P\right]$, where $a$ is $0, b$ is 107.075 , and $P$ is $\left[2 F_{c}^{2}+\operatorname{Max}\left(F_{0}^{2}, 0\right)\right] / 3$. |  |

distances are listed in Table 4 and bond valences in Table 5. CIF and structure factors are on deposit ${ }^{1}$.

## DESCRIPTION OF THE STRUCTURE

The most distinctive unit in the structure (Fig. 4) is an edgesharing octahedral tellurate dimer, $\left[\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right]^{8-}$. The octahedral dimers bond to Pb atoms, which in turn are linked via bonds to Cl atoms, $\mathrm{CO}_{3}$ triangles, and $\mathrm{H}_{2} \mathrm{O}$ molecules. The thorneite structure is unique and is not closely related to that of any other mineral.

The average Te-O bond length in thorneite, $1.934 \AA$, is typical of octahedrally coordinated $\mathrm{Te}^{6+}$ [e.g., frankhawthorneite, $1.939 \AA$ (Grice and Roberts 1995); jensenite, $1.936 \AA$ (Grice et al. 1996); leisingite, $1.922 \AA$ (Margison et al. 1997); khinite, $1.962 \AA$ (Cooper et al. 2008); ottoite, $1.942 \AA$ (Kampf et al. 2010a,); housleyite, 1.931 and $1.942 \AA$ (Kampf et al. 2010b); markcooperite, $1.95 \AA$ (Kampf et al. 2010c); timroseite, $1.933 \AA$
${ }^{1}$ Deposit item AM-10-045, CIF and structure factors. Deposit items are available two ways: For a paper copy contact the Business Office of the Mineralogical Society of America (see inside front cover of recent issue) for price information. For an electronic copy visit the MSA web site at http://www.minsocam.org, go to the American Mineralogist Contents, find the table of contents for the specific volume/issue wanted, and then click on the deposit link there.

TAbLE 4. Selected bond distances ( $\AA$ ) for thorneite

| Te-O3 |  | 1.889(10) |  | Pb1-O6 |  | 2.297(9) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Te-O4 |  | 1.901(9) |  | Pb1-O3 |  | 2.360(10) |
| Te-O6 |  | 1.930(9) |  | Pb1-O5 |  | 2.423(9) |
| Te-O5 |  | 1.931(9) |  | $\mathrm{Pb} 1-\mathrm{O} 2$ |  | 2.777(12) |
| Te-O7 |  | 1.974(9) |  | Pb1-O6 |  | 3.016(9) |
| Te-O7 |  | 1.979(8) |  | Pb1-01 |  | 3.155(1) |
| <Te-O> |  | 1.934 |  | Pb1-Cl |  | 3.456(7) |
|  |  |  |  | $\mathrm{Pb} 1-\mathrm{O} 2$ |  | 3.472(13) |
| C-O2(×2) |  | 1.269(15) |  | Pb1-O4 |  | 3.728(10) |
| C-O1 |  | 1.291(26) |  | Pb1-O3 |  | 3.806(10) |
| <C-O> |  | 1.276 |  | <Pb1-O,Cl> |  | 3.086 |
| Pb2-O4 |  | 2.439(9) |  | Pb3-O4 |  | 2.317(10) |
| Pb2-O5 |  | 2.490(9) |  | Pb3-O6 |  | 2.404(9) |
| Pb2-O3 |  | 2.554(10) |  | Pb3-O2 |  | 2.406(12) |
| $\mathrm{Pb} 2-\mathrm{O} 5$ |  | 2.661 (9) |  | Pb3-01 |  | 2.679(10) |
| Pb2-07 |  | 2.718(9) |  | Pb3-OW |  | 2.758(16) |
| Pb2-Cl |  | 2.968(6) |  | Pb3-Cl |  | 3.360(6) |
| Pb2-O3 |  | 3.096(11) |  | Pb3-Cl |  | 3.489(7) |
| Pb2-Cl |  | 3.599(7) |  | Pb3-O2 |  | 3.699(12) |
| Pb2-O7 |  | 3.808(8) |  | Pb3-07 |  | 3.788(9) |
| <Pb2-O,Cl> |  | 2.991 |  | <Pb3-O,Cl> |  | 3.068 |
| Hydrogen bonds ( $\mathrm{D}=$ donor, $\mathrm{A}=$ acceptor) |  |  |  |  |  |  |
| D-H Cd | d(D-H) | $\mathrm{d}(\mathrm{H}-\mathrm{A})$ | D-H-A | - d(D-A) | A | H-D-H |
| OW-H 0.9 | 0.92(3) | 2.69(21) | $114{ }^{\circ}$ | 3.175(6) | Cl | $104{ }^{\circ}$ |
| OW-H 0 | 0.92(3) | 2.71(21) | $128^{\circ}$ | 3.355(6) | Cl |  |

TABLE 3. Atomic coordinates and displacement parameters $\left(\AA^{2}\right)$ for thorneite

|  | $x$ | $y$ | $z$ | $U_{\text {eq }}$ | $U_{11}$ | $U_{22}$ | $U_{33}$ | $U_{23}$ | $U_{13}$ | $U_{12}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pb1 | 0.63753(3) | 0.09635(5) | $0.15303(7)$ | 0.0186(2) | 0.0139(3) | 0.0241(3) | 0.0165(3) | -0.0012(2) | -0.0005(2) | 0.0015(3) |
| Pb2 | $0.69515(3)$ | 0.41296(5) | $0.96922(8)$ | 0.0220(2) | 0.0218(4) | 0.0233(3) | 0.0195(3) | -0.0021(3) | 0.0007(3) | -0.0020(3) |
| Pb3 | 0.55264(3) | 0.24366(6) | 0.52254(8) | 0.0283(2) | 0.0177(4) | 0.0376(4) | 0.0279(4) | -0.0047(3) | 0.0006(3) | 0.0056(3) |
| Te | $0.78842(5)$ | 0.30010(8) | 0.36271 (12) | 0.0134(2) | 0.0114(5) | 0.0160(5) | 0.0128(5) | -0.0008(4) | 0.0024(4) | -0.0030(4) |
| C | 0.5 | 0.9822(20) | 0.25 | 0.025(6) | 0.028(16) | 0.017(13) | 0.025(13) | 0.0 | -0.010(11) | 0.0 |
| Cl* | 0.5884(3) | 0.3933(6) | 0.1682(8) | 0.057(3) | 0.032(4) | 0.080(6) | 0.059(5) | -0.002(3) | 0.009(3) | 0.006(3) |
| 01 | 0.5 | 0.0990(15) | 0.25 | 0.052(5) | 0.034(12) | 0.033(11) | 0.079(14) | 0.0 | -0.015(10) | 0.0 |
| 02 | 0.5487(6) | 0.9265(11) | 0.2183(17) | 0.044(3) | 0.022(7) | 0.049(8) | 0.063(9) | -0.009(7) | 0.013(6) | 0.001(6) |
| 03 | 0.7815(5) | 0.4526(9) | 0.2479(13) | 0.026(3) | 0.024(7) | 0.028(6) | 0.021(6) | 0.008(5) | -0.006(5) | -0.011(5) |
| 04 | 0.6451(5) | 0.2717(9) | $0.7367(12)$ | 0.020(2) | 0.016(6) | 0.028(6) | 0.016(5) | -0.012(5) | 0.003(5) | -0.008(5) |
| O5 | $0.7269(5)$ | 0.2338(8) | $0.1654(12)$ | 0.016(2) | 0.013(6) | 0.019(6) | 0.014(5) | 0.004(4) | -0.004(4) | -0.004(4) |
| O6 | $0.6435(4)$ | 0.1381(8) | $0.4534(12)$ | 0.014(2) | 0.008(5) | 0.025(6) | 0.010(5) | -0.002(4) | 0.003(4) | -0.001(4) |
| 07 | 0.7201(4) | 0.3460(8) | 0.4937 (12) | 0.012(2) | 0.011(6) | 0.013(5) | 0.014(5) | -0.003(4) | 0.008(4) | 0.002(4) |
| OW | 0.5 | 0.391(2) | 0.75 | 0.089(8) | 0.095(24) | 0.094(20) | 0.081(20) | 0.0 | 0.022(17) | 0.0 |
| H | 0.525(8) | 0.442(5) | 0.829 (19) | 0.08(8) |  |  |  |  |  |  |
| * Refined Cl:OH occupancy: 0.75(4):0.25(4). |  |  |  |  |  |  |  |  |  |  |

TABLE 5. Bond-valence analysis for thorneite

|  | 01 | O2 | O3 | O4 | O5 | O6 | 07 | OW | Cl | Sum |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Te |  |  | 1.079 | 1.044 | 0.963 | 0.965 | 0.857, 0.846 |  |  | 5.754 |
| C | 1.307 | $1.387 \times 2 \rightarrow$ |  |  |  |  |  |  |  | 4.080 |
| Pb1 | $0.088 \times 2 \downarrow$ | $0.190,0.046$ | 0.445, 0.023 | 0.027 | 0.391 | 0.506, 0.117 |  |  | 0.125 | 1.958 |
| Pb2 |  |  | 0.299, 0.099 | 0.379 | $0.341,0.241$ |  | 0.214, 0.023 |  | 0.339, 0.094 | 2.029 |
| Pb3 | $0.232 \times 2 \downarrow$ | 0.405, 0.029 |  | 0.486 |  | 0.407 | 0.024 | $0.197 \times 2 \downarrow$ | 0.152, 0.117 | 2.049 |
| H |  |  |  |  |  |  |  | 0.926 | 0.074 | 1.000 |
| $\mathrm{H}^{\prime}$ |  |  |  |  |  |  |  | 0.931 | 0.069 | 1.000 |
| Sum | 1.946 | 2.057 | 1.945 | 1.936 | 1.936 | 1.994 | 1.964 | 2.252 | 0.970 |  |

Notes: Values are expressed in valence units. Multiplicity is indicated by $\times \rightarrow \downarrow ; \mathrm{Pb}^{2+}-\mathrm{O}$ bond strengths from Krivovichev and Brown (2001); $\mathrm{Pb}^{2+}-\mathrm{Cl}$ bond strengths from Brese and O'Keeffe (1991); Te ${ }^{6+-} \mathrm{O}$ bond strengths and hydrogen-bond strengths based on $\mathrm{H} \cdots \mathrm{Cl} / \mathrm{H} \cdots \mathrm{O}$ bond lengths from Brown and Altermatt (1985); bond strength for Cl site assigned in accord with refined site occupancy $\left[\mathrm{Cl}_{0.75}(\mathrm{OH})_{0.25}\right]$.


Figure 4. Structure of thorneite viewed down c. Numbered spheres are Pb atoms ( $\mathrm{Pb} 1, \mathrm{~Pb} 2$ and Pb 3 ). Smallest spheres associated with OW are H atoms.
(Kampf et al. 2010d); and paratimroseite, $1.926 \AA$ (Kampf et al. 2010d); however, the $\mathrm{TeO}_{6}$ octahedron exhibits significant distortion, attributable to the repulsion between the $\mathrm{Te}^{6+}$ cations across the 07-O7 shared edge of the dimer. Te-O distances vary from 1.889 to $1.979 \AA$, $\mathrm{O}-\mathrm{Te}-\mathrm{O}$ angles vary from 77.3 to $100.9^{\circ}$ and $\mathrm{O}-\mathrm{O}$ distances vary from 2.469 to $2.923 \AA$. The $\mathrm{Te}-\mathrm{O} 7$ distances are the longest, the $\mathrm{O} 7-\mathrm{Te}-\mathrm{O} 7$ angle is the smallest and the 07-O7 edge is the shortest. The edge-sharing octahedral tellurate dimer, $\left[\mathrm{Te}_{2}^{6+} \mathrm{O}_{10}\right]^{8-}$, has not yet been found in any other mineral structure; however, it has been reported in synthetic compounds, cf. $\mathrm{K}_{4}\left[\mathrm{Te}_{2}^{6+} \mathrm{O}_{6}(\mathrm{OH})_{4}\right]\left(\mathrm{H}_{2} \mathrm{O}\right)_{7.3}$ (Lindqvist and Lundgren 1966), $\mathrm{Na}_{2} \mathrm{~K}_{4}\left[\mathrm{Te}_{2}^{6+} \mathrm{O}_{8}(\mathrm{OH})_{2}\right]\left(\mathrm{H}_{2} \mathrm{O}\right)_{14}$ (Lindqvist 1969), and $\mathrm{Na}_{4} \mathrm{TeO}_{5}$ (Untenecker and Hoppe 1987). Similar octahedral distortions related to $\mathrm{Te}^{6+}$ repulsions occur in all of these structures.

The Pb atoms have lopsided nine- and tenfold coordinations (Fig. 5) attributable to the $6 \mathrm{~s}^{2}$ lone-electron-pair effect typically exhibited by $\mathrm{Pb}^{2+}$ (e.g., Moore 1988; Cooper and Hawthorne 1994; Kharisun et al. 1997; Mills et al. 2009). In fact, the Pb atoms in the structures of all seven recently discovered new minerals from Otto Mountain exhibit this effect.

The two equivalent H atoms of the OW group form bifurcated H bonds to two equivalent Cl sites (shown in Fig. 4). The H bonds are relatively long ( 2.68 and $2.71 \AA$ ), leaving OW relatively


Figure 5. Coordinations of the $\mathrm{Pb} 1, \mathrm{~Pb} 2$, and Pb 3 atoms in thorneite. The lopsided distributions of bond lengths are attributable to the localization of the lone-pair electrons. Bond lengths are given in angstroms.
oversaturated. The 0.25 H atom associated with the 0.25 O at the Cl site may participate in an H bond to OW , but based upon the shortest Cl-OW bond length ( $3.175 \AA$ ) this would contribute a bond strength of about 0.005 valence units.

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## REFERENCES CITED

Altomare, A., Cascarano, G., Giacovazzo, C., Guagliardi, A., Burla, M.C., Polidori, G., and Camalli, M. (1994) SIR92-A program for automatic solution of crystal structures by direct methods. Journal of Applied Crystallography, 27, 435.
Brese, N.E. and O'Keeffe, M. (1991) Bond-valence parameters for solids. Acta Crystallographica, B47, 192-197.
Brown, I.D. and Altermatt, D. (1985) Bond-valence parameters from a systematic analysis of the inorganic crystal structure database. Acta Crystallographica, B41, 244-247.
Cooper, M. and Hawthorne, F.C. (1994) The crystal structure of wherryite, $\mathrm{Pb}_{7} \mathrm{Cu}_{2}\left(\mathrm{SO}_{4}\right)_{4}\left(\mathrm{SiO}_{4}\right)_{2}(\mathrm{OH})_{2}$, a mixed sulfate-silicate with $\left[{ }^{[6]} \mathrm{M}\left(\mathrm{TO}_{4}\right)_{2} \Phi\right]$ chains. Canadian Mineralogist, 32, 373-380
Cooper, M.A., Hawthorne, F.C., and Back, M.E. (2008) The crystal structure of khinite and polytypism in khinite and parakhinite. Mineralogical Magazine, 72, 763-770.
Grice, J.D. and Roberts, A.C. (1995) Frankhawthorneite, a unique HCP framework structure of a cupric tellurate. Canadian Mineralogist, 33, 649-653.
Grice, J.D., Groat, L.A., and Roberts, A.C. (1996) Jensenite, a cupric tellurate framework structure with two coordinations of copper. Canadian Mineralogist, 34, 55-59.
Holland, T.J.B. and Redfern, S.A.T. (1997) Unit cell refinement from powder diffraction data: the use of regression diagnostics. Mineralogical Magazine, 61, 65-77.
Kampf, A.R., Rossman, G.R., and Housley, R.M. (2009) Plumbophyllite, a new species from the Blue Bell claims near Baker, San Bernardino County, California. American Mineralogist, 94, 1198-1204
Kampf, A.R., Housley, R.M., Mills, S.J., Marty, J., and Thorne, B. (2010a) Lead-tellurium oxysalts from Otto Mountain near Baker, California: I. Ottoite, $\mathrm{Pb}_{2} \mathrm{TeO}_{5}$, a new mineral with chains of tellurate octahedra. American Mineralogist, 95, 1329-1336.
Kampf, A.R., Marty, J., and Thorne, B. (2010b) Lead-tellurium oxysalts from Otto Mountain near Baker, California: II. Housleyite, $\mathrm{Pb}_{6} \mathrm{CuTe}_{4} \mathrm{O}_{18}(\mathrm{OH})_{2}$, a new mineral with $\mathrm{Cu}-\mathrm{Te}$ octahedral sheets. American Mineralogist, 95, 1337-1343.
Kampf, A.R., Mills, S.J., Housley, R.M., Marty, J., and Thorne, B. (2010c) Lead-tellurium oxysalts from Otto Mountain near Baker, California: IV,

Markcooperite, $\mathrm{Pb}_{2}\left(\mathrm{UO}_{2}\right) \mathrm{Te}^{6+} \mathrm{O}_{6}$, the first natural uranyl tellurate. American Mineralogist, 95, 1554-1559.
(2010d) Lead-tellurium oxysalts from Otto Mountain near Baker, California: V. Timroseite, $\mathrm{Pb}_{2} \mathrm{Cu}_{5}^{2+}\left(\mathrm{Te}^{6+} \mathrm{O}_{6}\right)_{2}(\mathrm{OH})_{2}$, and paratimroseite, $\mathrm{Pb}_{2} \mathrm{Cu}_{4}^{2+}\left(\mathrm{Te}^{6+} \mathrm{O}_{6}\right)_{2}\left(\mathrm{H}_{2} \mathrm{O}\right)_{2}$, two new tellurates with $\mathrm{Te}-\mathrm{Cu}$ polyhedral sheets. American Mineralogist, 95, 1560-1568.
(2010e) Lead-tellurium oxysalts from Otto Mountain near Baker, California: VI. Telluroperite, $\mathrm{Pb}_{3} \mathrm{Te}^{4+} \mathrm{O}_{4} \mathrm{Cl}_{2}$, the Te analog of perite and nadorite. American Mineralogist, 95, 1569-1573.
Kharisun, Taylor, M.R., Bevan, D.J.M., Rae, A.D., and Pring, A. (1997) The crystal structure of mawbyite, $\mathrm{PbFe}_{2}\left(\mathrm{AsO}_{4}\right)_{2}(\mathrm{OH})_{2}$. Mineralogical Magazine, 61, 685-691.
Kraus, W. and Nolze, G. (1996) POWDER CELL-A program for the representation and manipulation of crystal structures and calculation of the resulting X-ray powder patterns. Journal of Applied Crystallography, 29, 301-303.
Krivovichev, S.V. and Brown, I.D. (2001) Are the compressive effects of encapsulation an artifact of the bond valence parameters? Zeitschrift für Kristallographie, 216, 245-247.
Lindqvist, O. (1969) The crystal structure of the tellurate $\mathrm{Na}_{2} \mathrm{~K}_{4}\left[\mathrm{Te}_{2} \mathrm{O}_{8}(\mathrm{OH})_{2}\right]$ $\left(\mathrm{H}_{2} \mathrm{O}\right)_{14}$. Acta Chemica Scandinavica, 23, 3062-3070
Lindqvist, O. and Lundgren, G. (1966) The crystal structure of potassium tellurate (VI). Acta Chemica Scandinavica, 20, 2138-2155.

Mandarino, J.A. (1981) The Gladstone-Dale relationship. IV. The compatibility concept and its application. Canadian Mineralogist, 19, 441-450.
Margison, S.M., Grice, J.D., and Groat, L.A. (1997) The crystal structure of leisingite, $\left(\mathrm{Cu}^{2+}, \mathrm{Mg}, \mathrm{Zn}\right)_{2}(\mathrm{Mg}, \mathrm{Fe}) \mathrm{Te}^{6+} \mathrm{O}_{6} \cdot 6 \mathrm{H}_{2} \mathrm{O}$. Canadian Mineralogist, 35, 759-763.
Mills, S.J., Kampf, A.R., Raudsepp, M., and Christy, A.G. (2009) The crystal structure of Ga-rich plumbogummite from Tsumeb, Namibia. Mineralogical Magazine, 73, 837-845
Moore, P.B. (1988) The joesmithite enigma: Note on the $6 s^{2} \mathrm{~Pb}^{2+}$ lone pair. American Mineralogist, 73, 843-844.
Sheldrick, G.M. (2008) A short history of SHELX. Acta Crystallographica, A64, 112-122.
Untenecker, H. and Hoppe, R. (1987) Ein Neues Oxotellurat, $\mathrm{Na}_{4} \mathrm{TeO}_{5}$, und eine Revision der Structur von $\mathrm{Li}_{4} \mathrm{TeO}_{5}$. Journal of the Less-Common Metals, 132, 79-92.

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